On the numerical implementation of a fully coupled gravity-wave ray tracer in atmospheric models

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- GWs are not fully resolved by GCMs and NWP models ⇒ parametrization ⇒ (Wentzel-Kramers-Brillouin) WKB theory
- Currently used parametrizations: equilibrium profile of the wave field, steady state background wind, fully linear theory ⇒ the resolved flow (mean flow) does not know about the GWs propagating through it ⇒ the GW momentum is deposited instantenously to the mean flow at height of wave breaking
- Proposal for improvement: weakly nonlinear coupling between the waves and the resolved flow ⇒ continuous interaction between the wave and the mean flow during the propagation

Wave resolving equations (2-D Euler equations, no rotation):

$$\frac{Du}{Dt} + c_p \theta \frac{\partial \pi}{\partial x} = 0$$
$$\frac{Dw}{Dt} + c_p \theta \frac{\partial \pi}{\partial z} + g = 0$$
$$\frac{D\theta}{Dt} = 0$$
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$$\frac{D\pi}{Dt} + \frac{\kappa}{1 - \kappa} \pi \left(\frac{\partial u}{\partial x} + \frac{\partial w}{\partial z}\right) = 0$$

with Exner pressure 7 Pot temperature 6

$$\begin{split} \frac{D}{Dt} &= \frac{\partial}{\partial t} + u \frac{\partial}{\partial x} + w \frac{\partial}{\partial z} \\ \pi &= (p/p_0)^{\kappa} \\ \theta &= T(p_0/p)^{\kappa} = T/\pi \\ \kappa &= R/c_p \end{split}$$

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$$f_w(x, z, t) = \operatorname{Re} F_w(Z, T) e^{i \left[kx + \frac{\phi(Z, T)}{\epsilon}\right]}$$

with $Z = \epsilon z, T = \epsilon t, \ m = \partial \phi / \partial Z$ and $\omega = -\partial \phi / \partial T$

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- Scaling for the gravity waves: $\epsilon = L_w/H_{ heta} << 1$: weak stratification

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• The coupled system (*Achatz et al.*, 2010, *Muraschko et al.*, 2015): <u>Wave field</u> <u>Mean flow</u>

$$\begin{aligned} \frac{\mathrm{d}_{g}z}{\mathrm{d}t} &= \quad \mp \frac{Nkm}{(k^{2}+m^{2})^{3/2}} \equiv c_{gz} & \frac{\partial}{\partial t} u_{b} \\ \frac{\partial}{\partial t} &= \quad -\frac{1}{\overline{\rho}} \frac{\partial}{\partial z} \left[\frac{\overline{\rho}}{2} \operatorname{Re} \left(U_{w} W_{w}^{*} \right) \right] \\ \frac{\mathrm{d}_{g}m}{\mathrm{d}t} &= \quad \mp \frac{k}{(k^{2}+m^{2})^{1/2}} \frac{\mathrm{d}N}{\mathrm{d}z} - k \frac{\mathrm{d}^{\mathbf{u}}u_{b}}{\mathrm{d}z} \equiv \dot{m} &= \quad -\frac{1}{\overline{\rho}} \frac{\partial}{\partial z} \left(kc_{gz} \mathbf{\mathcal{A}} \right) \\ \frac{\mathrm{d}_{g}\mathbf{\mathcal{A}}}{\mathrm{d}t} &= \quad -\mathbf{\mathcal{A}} \frac{\partial c_{gz}}{\partial z} \quad \left(\frac{\mathrm{d}_{g}}{\mathrm{d}t} = \frac{\partial}{\partial t} + c_{gz} \frac{\partial}{\partial z} \right) \quad \overline{\rho} &= \quad \rho(z) = \rho_{0} e^{-z/H} \quad H = \frac{RT_{0}}{g} \end{aligned}$$

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• Problem: if rays crossing \Rightarrow caustics: several m at same height $z \Rightarrow$ e.g. c_{gz} multivalued BUT! $\mathcal{A} = \mathcal{A}(z, t) \Rightarrow$ wave action conservation ill-defined \Rightarrow numerical problems

• Solution: extension of the model to a 2D phase space (z, m)



Hertzog et al., 2002, Muraschko et al., 2015

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• "Slicing up" the wave action density to several m intervals \Rightarrow phase-space wave action density:

$$\mathcal{N}(z,m,t) = \int\limits_{R} \mathcal{A}_{\alpha}(z,t) \delta[m-m_{\alpha}(z,t)] d\alpha$$



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- multiple m values allowed at each location z → spectral (non-monochromatic) treatment → no caustics problems
- extension to Muraschko et al., 2015: isothermal background with variable density $\overline{\rho} = \rho(z) = \rho_0 e^{-z/H}$ $H = \frac{RT_0}{g}$

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- saturation occurs if static instability sets in (*Lindzen 1981*), i.e. $\partial \theta_w / \partial z + d\overline{\theta} / dz < 0$ or, after an additional multiplication by $g/\overline{\theta}$

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 this monochromatic criterion can be generalized for a spectrum by taking its integral over the vertical wavenumber:

$$\int_{-\infty}^{\infty} (m|B_w|)^2 dm = \int_{-\infty}^{\infty} m^2 f(\mathcal{N}) dm > \alpha^2 N^4$$

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• If the saturation criterion is fulfilled the wave action density \mathcal{N} is reset layer-wise to a value that sets back stability

Methodology

- LES: fully non-linear wave resolving reference (PincFloit, Rieper et al., 2013)
- WKB-eu: Eulerian WKB model
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Cases

- refraction by a jet: U_{jet} weak \Rightarrow m only slightly modified
- reflection from a jet: $U_{jet} \geq rac{N}{k} \left(1 rac{k}{\sqrt{k^2 + m^2}}\right) \Rightarrow m$ and c_{gz} changes sign
- static instability: amplitude growing until static instability
- modulational instability: $|m| \approx |k| \Rightarrow$ wave packet is shrinking, its amplitude growing \rightarrow static instability
- Critical layer: $U_{jet} \approx -c_p \Rightarrow m$ grows to infinity, wavepacket collapses

Induced wind, refraction by a jet $(\lambda_x = 10km, \lambda_z = 1km)$



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Wave energy, reflection from a jet $(\lambda_x = 10km, \lambda_z = 1km)$



Wave energy, partial reflection from a jet ($\lambda_x = 6km, \lambda_z = 3km$)



Static instability ($\lambda_x = \lambda_z = 1km$)



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- The Lagrangian WKB model has been supplemented with a wave breaking parametrization, which can slightly improve the fit to LES.
- However! the WKB model without the breaking parametrization works surprisingly well → the most important part of the wave - background interaction is the weakly-nonlinear coupling and NOT wave breaking or wave-wave interactions
- There is a good reason to try out fully coupled transient approaches in GW parametrizations and go beyond the non-acceleration theorem applied in most current schemes

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Questions?

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Experiment	Wavepacket	Background	Domain size	Resolution
REFR	Cosine shape	non-Boussinesq	WKB Euler:	WKB Euler:
Refraction	$\lambda_x = 10 km, \lambda_z = 1 km$	T = 300K	$L_z = 40 km$	nz = 400, nm = 70
by a jet	$k=2\pi/\lambda_{\rm x}, m=2\pi/\lambda_{\rm z}$	$N \approx 0.018$	$m \in [0.001, 0.008]$	$dz \approx 100m, dm = 10^{-4}s^{-1}$
	$z_0=10 km, \Delta_{wp}=10 km$	$u_0 = 5m/s$	WKB Lagrange:	WKB Lagrange:
	$branch=-1, a_0=0.1$	$z_{\alpha} = 25km$	$L_z = 40 km$	$nz = 400, dz_{smooth} \approx 600m$
		$\Delta_u = 10 km$		$dz \approx 100m, n_{ray} = 4000$
			LES:	LES:
			$L_{\rm z}=40km, L_{\rm x}=10km$	nz = 1280, nx = 32
				$dz \approx 31m, dx = 310m$
REFL	Cosine shape	non-Boussinesq	WKB Euler:	WKB Euler:
Reflection	$\lambda_x = 10 km, \lambda_z = 1 km$	T = 300K	$L_z = 40 km$	nz = 400, nm = 180
from a jet	$k=2\pi/\lambda_{\rm x}, m=2\pi/\lambda_{\rm c}$	$N \approx 0.018$	$m \in [-0.01, 0.008]$	$dz \approx 100m, dm = 10^{-4}s^{-1}$
	$z_0=10 km, \Delta_{wp}=10 km$	$u_0=40m/s$	WKB Lagrange:	WKB Lagrange:
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		$\Delta_u = 10 km$		$dz\approx 100m, n_{ray}=4000$
			LES:	LES:
			$L_{\rm z}=40km, L_{\rm x}=10km$	nz = 2500, nx = 64
				$dz \approx 16m, dx = 156m$
PREFL	Cosine shape	non-Boussinesq	WKB Lagrange:	WKB Lagrange:
Partial	$\lambda_x = 6km, \lambda_z = 3km$	T = 300K	$L_z = 50 km$	$nz = 166, dz_{smooth} \approx 1800m$
Reflection	$k=2\pi/\lambda_{\rm x}, m=2\pi/\lambda_{\rm z}$	$N \approx 0.018$		$dz \approx 300m, n_{ray} = 4320$
from a jet	$z_0=10 km, \Delta_{wp}=10 km$	$u_0=9,75m/s$	LES:	LES:
	$branch=-1, a_0=0.1$	$z_{\alpha} = 25km$	$L_z = 50 km, L_x = 6 km$	nz = 538, nx = 32
		$\Delta_u = 10 km$		$dz \approx 93m, dx = 187m$

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STIH	Gaussian shape	non-Boussinesq	WKB Lagrange:	WKB Lagrange:	
Static	$\lambda_x = 30 km, \ \lambda_z = 3 km$	T = 300K	$L_z = 80 km$	$nz=266, dz_{amooth}\approx 1800m$	
Instability	$k = 2\pi/\lambda_x, m = 2\pi/\lambda_z$	$N \approx 0.018$		$dz \approx 300m, n_{ray} = 4320$	
Hydrostatic	$z_0=10 km, \Delta_{wp}=25 km$		LES:	LES:	
Wavepacket	$branch = -1, a_0 = 0.5$		$L_z = 80 km, L_x = 30 km$	nz = 854, nx = 32	
				$dz \approx 94m, dx \approx 940m$	
STINH	Gaussian shape	non-Boussinesq	WKB Lagrange:	WKB Lagrange:	
Static	$\lambda_x = 1km, \lambda_z = 1km$	T = 300K	$L_z = 30km$	$nz=300, dz_{anooth}\approx 600m$	
Instability	$k = 2\pi/\lambda_x, m = 2\pi/\lambda_z$	$N \approx 0.018$		$dz \approx 100m, n_{ray} = 4000$	
Non-hydrostatic	$z_0 = 10 km, \Delta_{wp} = 10 km$		LES:	LES:	
Wavepacket	$branch = -1, a_0 = 0.9$		$L_z = 30km, L_x = 1km$	nz = 960, nx = 32	
				$dz \approx 31m, dx \approx 310m$	
МІ	Cosine shape	non-Boussinesq	WKB Lagrange:	WKB Lagrange:	
Modulational	$\lambda_x = 1km, \lambda_z = 1km$	T = 300K	$L_z = 60 km$	$nz=600, dz_{amooth}\approx 600m$	
Instability	$k=2\pi/\lambda_{\rm x}, m=2\pi/\lambda_{\rm z}$	$N \approx 0.018$		$dz \approx 100m, n_{ray} = 4000$	
	$z_0=10 km, \Delta_{wp}=20 km$		LES:	LES:	
	$branch = -1, a_0 = 0.1$		$L_{\rm c}=60km, L_{\rm x}=1km$	nz = 1920, nx = 32	
				$dz \approx 31m, dx \approx 310m$	
CL	Cosine shape	non-Boussinesq	WKB Lagrange:	WKB Lagrange:	
Critical	$\lambda_x = 10km, \lambda_z = 1km$	T = 300K	$L_z = 30 km$	$nz = 300, dz_{smooth} \approx 600m$	
Layer	$k = 2\pi/\lambda_x, m = 2\pi/\lambda_z$	$N \approx 0.018$		$dz\approx 100, n_{my}=4000$	
	$z_0 = 10 km, \Delta_{wp} = 10 km$	$u_0 = -11m/s$	LES:	LES:	
	$branch = -1, a_0 = 0.1$	$z_u = 25 km$	$L_z = 30 km, L_x = 10 km$	nz = 960, nx = 32	
		$\Delta_{a} = 10 km$		$dz \approx 31m, dx \approx 310m$	
					A = 1

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