

Gravity wave emission and propagation in the differentially heated rotating annulus experiment

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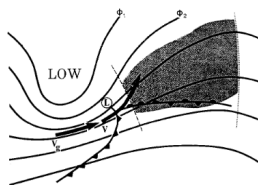
2016 SPARC Gravity Wave Symposium

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Introduction and Motivation

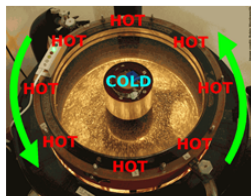
- Orographic and convective sources of inertia-gravity waves (IGW) well understood (Fritts (2003))
- Spontaneously emitted GWs active field of research (Cámara and Lott (2015), ...)
- Observation identify increased IGW activity in jet exit regions (e.g. Plougonven and Zhang (2014))
- GW emission embedded in various atmospheric processes
- Need for controllable, repeatable and simplified laboratory experiments: **rotating annulus**



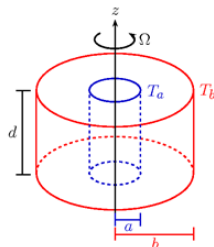
Koch and OHandley (1997)

Introduction and Motivation

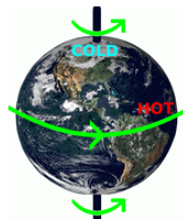
- Differentially heated rotating annulus experiment:



Laboratory experiment (BTU Cottbus)



schematic view

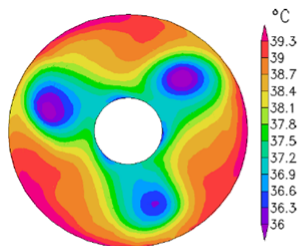


analogous to Earth's atmosphere

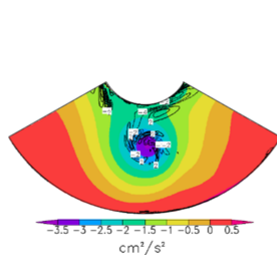
- Finite volume code (cylFloit) to simulate experiment using Boussinesq approximation (Borchert et al. (2015))

Introduction and Motivation

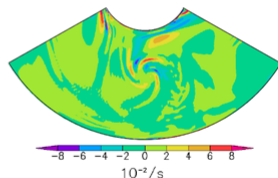
- Annulus simulations show clear GW activity (Borchert et al. (2014))



Development of baroclinic waves



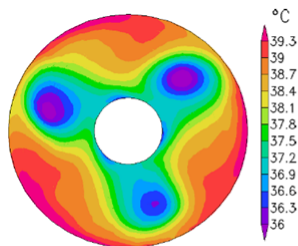
Pressure and horizontal divergence



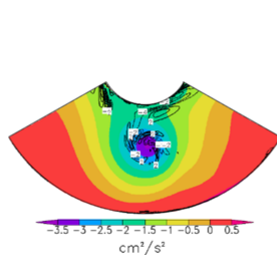
$$\delta = \nabla_h \cdot \vec{u}$$

Introduction and Motivation

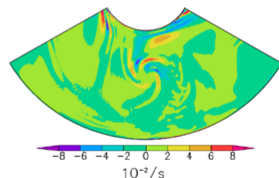
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Development of baroclinic waves



Pressure and horizontal divergence



$$\delta = \nabla_h \cdot \vec{u}$$

- Further understanding of GW source processes: tangent linear analysis (see Snyder et al. (2009), Wang and Zhang (2010))

⇒ **Is there some internal forcing of GWs by the balanced flow?**

Source mechanism of gravity wave emission

Decomposition of flow into **geostrophic** and **ageostrophic** part

$$\mathbf{v} = \mathbf{u}_g + \mathbf{v}_a$$

$$B = B_g + B_a$$

$$p = p_g + p_a$$

Source mechanism of gravity wave emission

Decomposition of flow into **geostrophic** and **ageostrophic** part

$$\begin{aligned} f \mathbf{e}_z \times \mathbf{u}_g + \nabla_h p_g &= 0 \\ B_g - \frac{\partial p_g}{\partial z} &= 0 \\ \mathbf{v} &= \mathbf{u}_g + \mathbf{v}_a \\ B &= B_g + B_a \\ p &= p_g + p_a \\ \Pi_g = \zeta + \frac{f}{N^2} \frac{\partial B}{\partial z} &= \frac{1}{f} \left(\nabla_h^2 + \frac{f^2}{N^2} \frac{\partial^2}{\partial z^2} \right) p_g \\ \zeta_a + \frac{f}{N^2} \frac{\partial B_a}{\partial z} &= 0 \\ \delta = \nabla_h \cdot \mathbf{u} = \delta_a &= \nabla_h \cdot \mathbf{u}_a \end{aligned}$$

Source mechanism of gravity wave emission

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... \Rightarrow Geostrophic forcing of ageostrophic flow

$$\frac{D\delta_a}{Dt} = -\frac{\partial B_a}{\partial z} + \frac{\partial^2 p_{aa}}{\partial z^2} + \frac{\partial \mathbf{v}}{\partial z} \cdot \mathbf{w}_a - \frac{\partial^2}{\partial z^2} \nabla^{-2} (\nabla \mathbf{u}_g \cdot \nabla \mathbf{u}_g)$$

Linearisation of unbalanced flow about balanced flow: Principle

- Decomposition into balanced (large) and unbalanced (small) part

$$x = \tilde{x} + x' \text{ with } |x'| \ll |\tilde{x}| \text{ (Unbalanced part} \equiv \text{gravity waves)}$$

- Tangent linear evolution of x'

$$\frac{\partial x'}{\partial t} = L(\tilde{x})x' + F(\tilde{x})$$

with a linear operator $L(\tilde{x})$ and a balanced forcing term $F(\tilde{x})$

- Unbalanced component is integrated separately within each time step
- Balanced part serves as background of the tangent linear model

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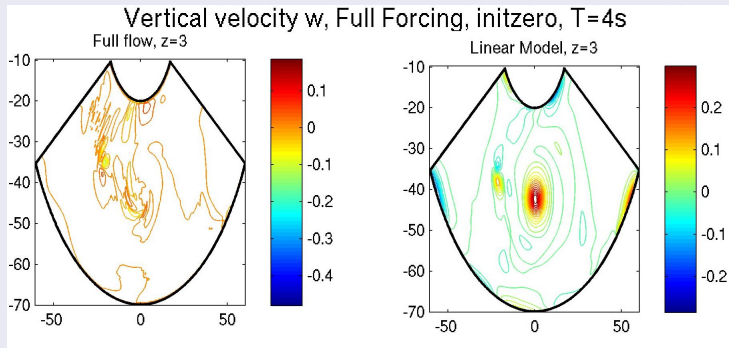
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-
- **Balanced forcing leading contributor to the gravity wave activity?**

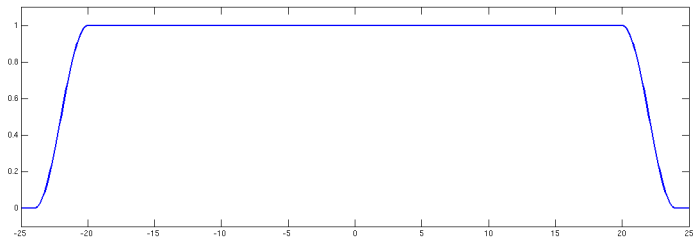
Problem: Instabilities at side walls

- linear model diverges after about 4-5 s of integration time
- Exponential growth rate at outer (and inner) side walls



Suppress growth at side wall: Multiplication with window function

$$f(x) = \begin{cases} 1, & |x| \leq \beta L_y \\ \frac{1}{2} \left\{ 1 + \cos \left[\frac{\pi(|x| - \beta L_y)}{L_s} \right] \right\}, & \beta L_y < |x| \leq [\beta + \gamma(1 - \beta)] L_y \\ 0, & \text{else} \end{cases}$$

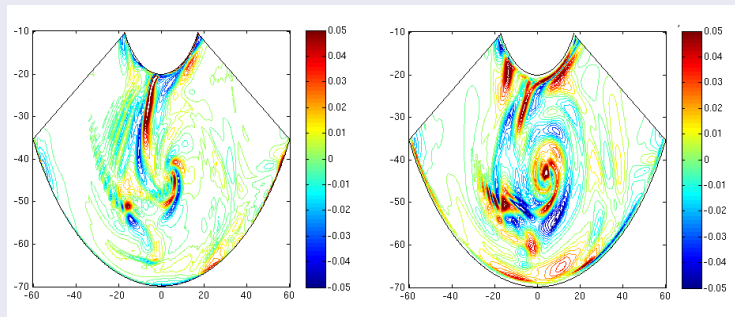


Comparison of horizontal divergence

- Initialising linear model with zero unbalanced part
- Full forcing
- $T=22$ s ($\Omega = 0.08$ rad/s, rotation period ~ 79 s)

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nonlinear model

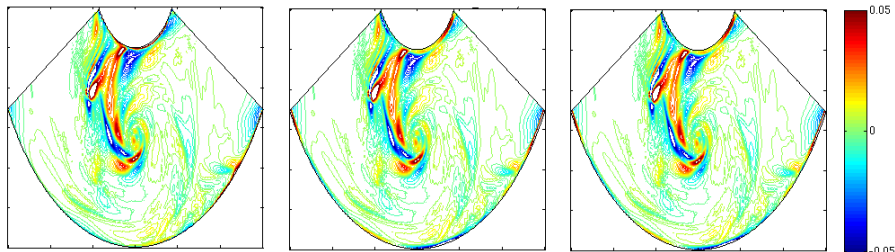
linear model

Comparison of horizontal divergence

- Initialising linear model with $x'_{init} = x - \tilde{x}$
- $T=0$ s

Comparison of horizontal divergence

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nonlinear model

linear with forcing

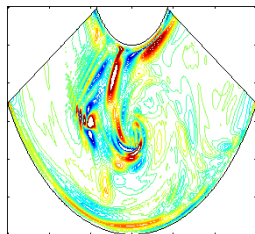
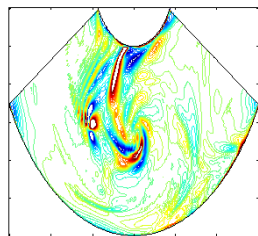
linear without forcing

Comparison of horizontal divergence

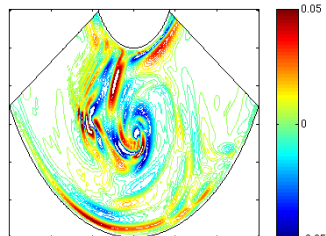
- Initialising linear model with $x'_{init} = x - \tilde{x}$
- $T=10$ s

Comparison of horizontal divergence

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linear with forcing



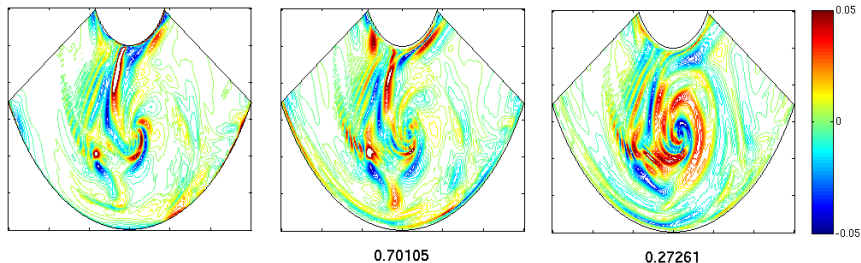
linear without forcing

Comparison of horizontal divergence

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nonlinear model

linear with forcing

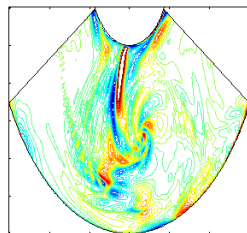
linear without forcing

Comparison of horizontal divergence

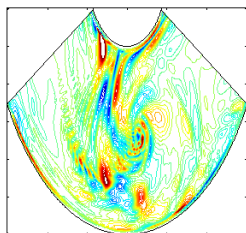
- Initialising linear model with $x'_{init} = x - \tilde{x}$
- $T=30$ s

Comparison of horizontal divergence

- Initialising linear model with $x'_{init} = x - \tilde{x}$
- $T=30$ s

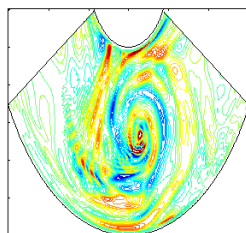


nonlinear model



0.54807

linear with forcing



0.076933

linear without forcing

Conclusion:

- Increased GW activity within the baroclinic wave and close to the inner cylinder wall
- Tangent linear analysis to gain further understanding of the GW source mechanism
- Window function to suppress growth rate at side walls
- Significant internal forcing of GW by the balanced flow

Outlook:

- Extract balanced part of horizontal divergence: omega equation
- Characterizing wave properties (\vec{k} , A , ..)

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- Borchert, S., U. Achatz, S. Remmler, S. Hickel, U. Harlander, M. Vincze, K. D. Alexandrov, F. Rieper, T. Heppelmann, and S. I. Dolaptchiev (2015, 01). Finite-volume models with implicit subgrid-scale parameterization for the differentially heated rotating annulus. *Meteorologische Zeitschrift* 23(6), 561–580.
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Tangent linear annulus equations

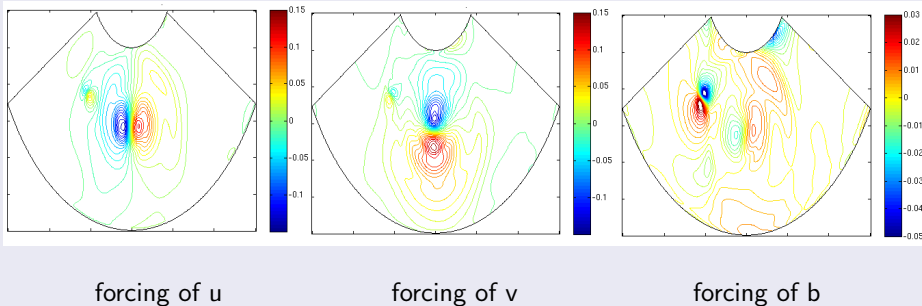
$$\frac{dB_a}{dt} = -N^2 w_a - \left(\frac{dB_g}{dt} \right)_a - \left\{ \left(\frac{dB_g}{dt} \right)_g \right\}$$

$$\frac{d\mathbf{u}_a}{dt} = -f \mathbf{e}_z \times \mathbf{u}_a - \nabla_h \tilde{p}_{aa} - \left(\frac{d\mathbf{u}_g}{dt} \right)_a - \left\{ \nabla_h \tilde{p}_{ag} + \left(\frac{d\mathbf{u}_g}{dt} \right)_g \right\}$$

$$\frac{dw_a}{dt} = B_a - \frac{\partial \tilde{p}_{aa}}{\partial z} - \left\{ \frac{\partial \tilde{p}_{ag}}{\partial z} \right\}$$

Forcing terms

- **Large scale** balanced forcing leading contributor to gravity wave activity
- $T=0$ s



Balanced part of horizontal divergence: Omega equation

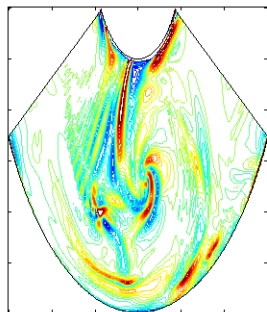
- Total horizontal divergence includes balanced part

$$\delta_{total} = \delta_{unbal} + \delta_{bal}$$

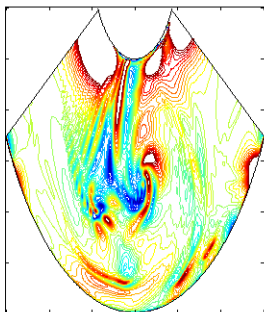
- Subtract balanced part using omega equation (Hoskins et al. (1978), Danioux et al. (2012))

$$\begin{aligned}\Rightarrow \delta_{unbal} &= \delta_{total} - \delta_{bal} \\ &= \nabla_h \cdot \mathbf{u}_a - \delta_{bal}, \text{ with } \delta_{bal} = -\frac{\partial w_{bal}}{\partial z} \\ \nabla_h^2 w_{bal} &= -\frac{2}{N^2} \nabla_h \cdot \mathbf{Q}, \text{ with } \mathbf{Q} = \nabla_h \mathbf{u}_g \cdot \nabla_h b_g\end{aligned}$$

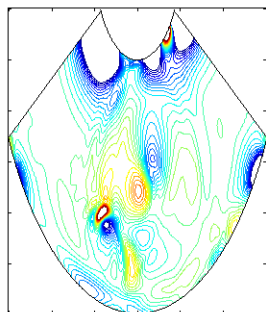
Balanced part of horizontal divergence: Omega equation



δ_{total}



δ_{unbal}



δ_{bal}

